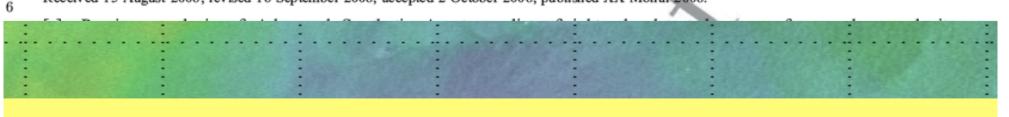


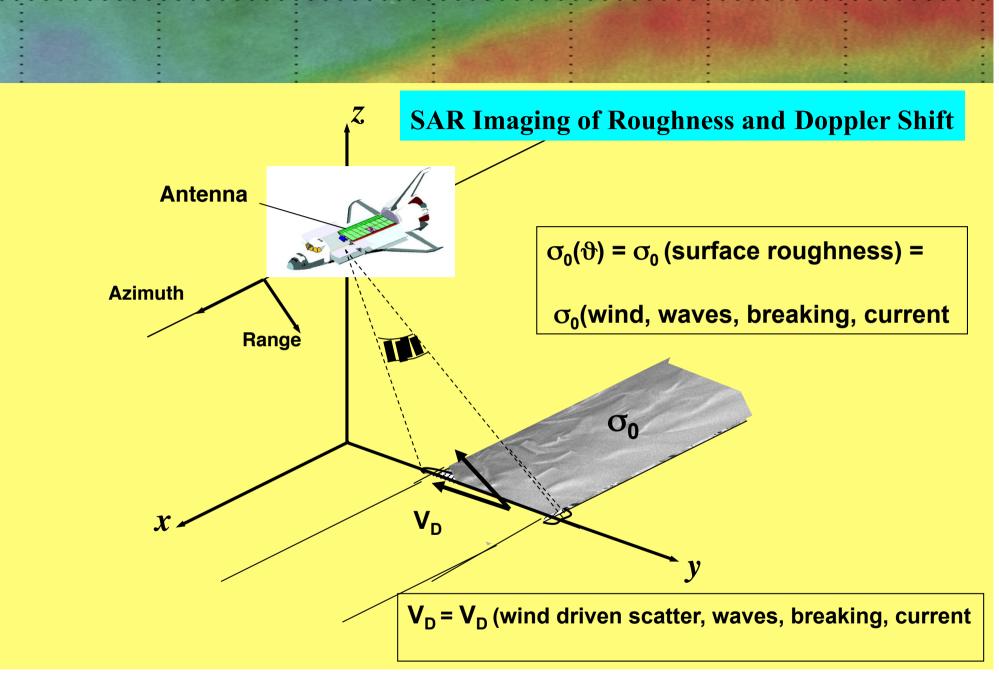
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- 2 Direct ocean surface velocity measurements from space: Improved
- **3 quantitative interpretation of Envisat ASAR observations**
- 4 J. A. Johannessen,^{1,2} B. Chapron,³ F. Collard,⁴ V. Kudryavtsev,^{1,5,6} A. Mouche,⁴
- 5 D. Akimov,⁵ and K.-F. Dagestad¹

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Estimation of Doppler anomaly

- Anomaly = measured predicted
- Compensated non-geophysical sources of anomaly :
 - Antenna misspointing
 - Instrumental bias
 - Doppler estimator bias caused by azimuthal variation of backscatter (artificial correlation between doppler and sigma0).
- Retrieval accuracy about 5 Hz which is about 0.2 m/s



Estimation of Doppler anomaly

Anomaly (shift) = measured – predicted

$$\frac{\pi f_D}{k_R} = -\frac{(u \sin \theta - w \cos \theta)\sigma_0(\theta + \Delta \theta)}{\sigma_0(\theta + \Delta \theta)}$$

$$f_0: \text{ Doppler shift}$$

$$K_R: \text{ Radar Wave number}$$

$$\theta: \text{ Incidence angle}$$

$$U: \text{ horizontal velocity}$$

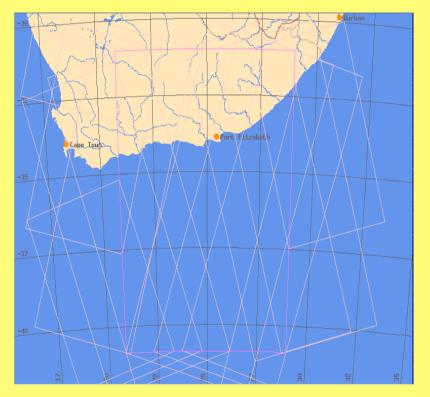
$$W: \text{ vertical velocity}$$

$$\sigma_0: \text{ radar cross section}$$

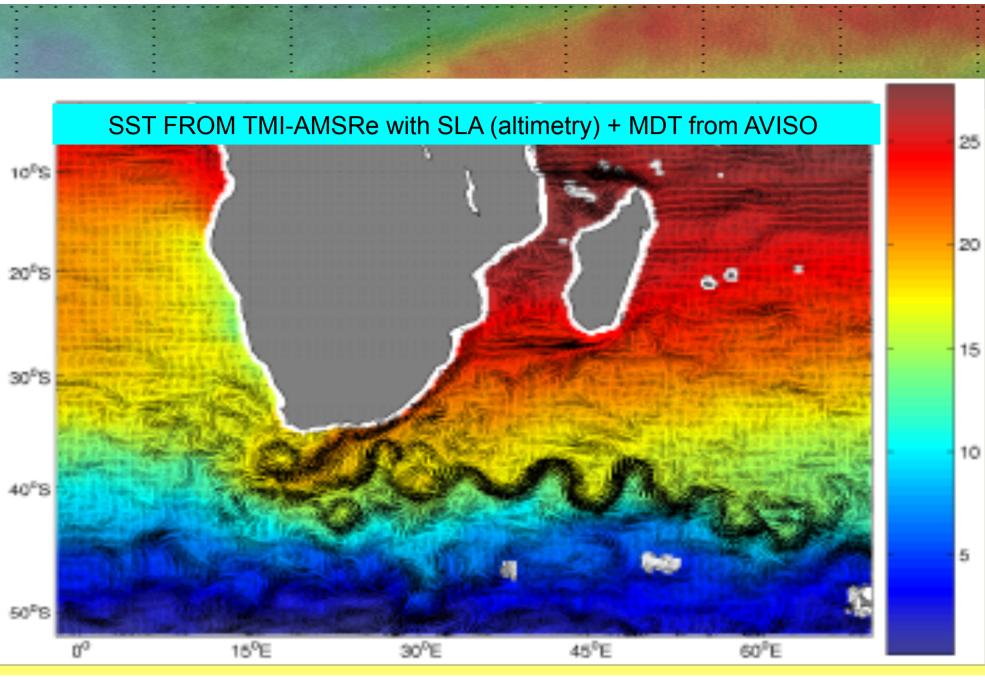


Study Area

Agulhas Current is excellent strong and intense current to be used as natural laboratories to explore satellite synergy (SAR, IR and OC) and develop new methods for quantitative estimates of current information and surface dynamics.

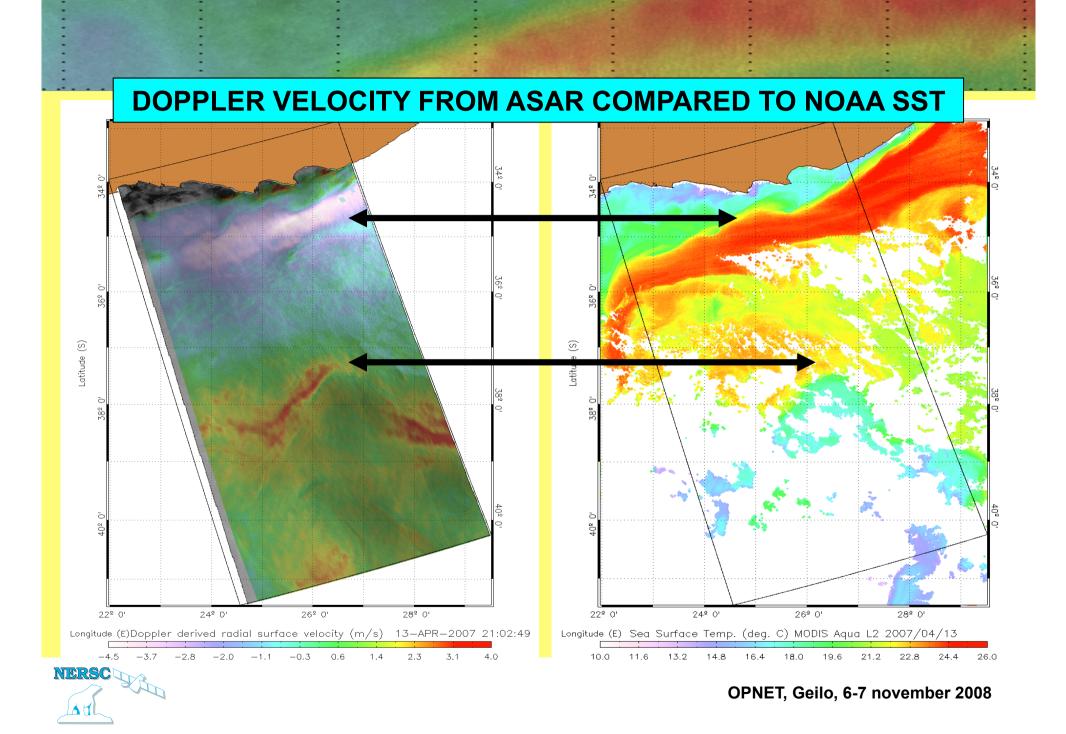








OPNET, Geilo, 6-7 november 2008



How is the wind effecting the results

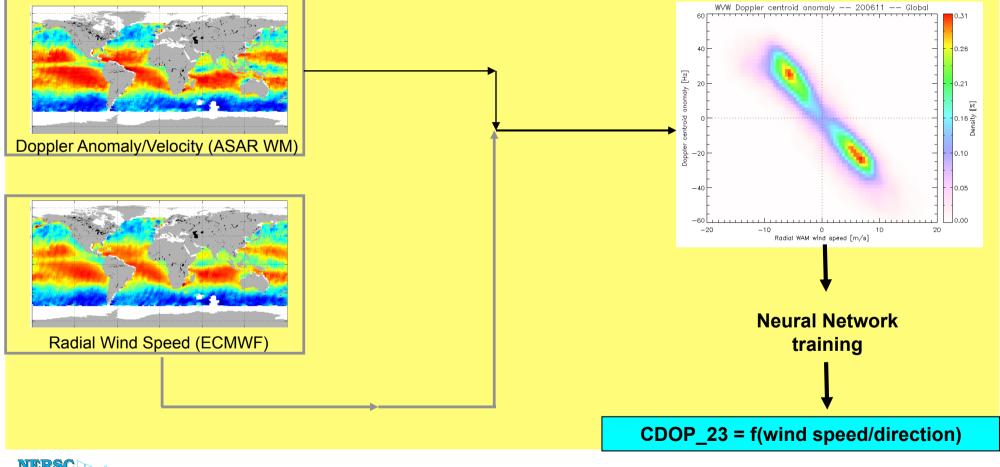
- the wind speed dependency of the Doppler signal for each month is evaluated using the global collocated (ASAR and ECMWF winds) data set,
- this allows the establishment of an empirical CDOP model relating observed Doppler anomalies to wind speed at different incidence angles, polarization, azimuth directions,
- in turn the removal of the wind contribution to the Doppler anomalies is possible

Following this the simulations of NRCS and Doppler anomalies is then consistently assessed and compared to the observations and empirical retrievals.



RESULTS

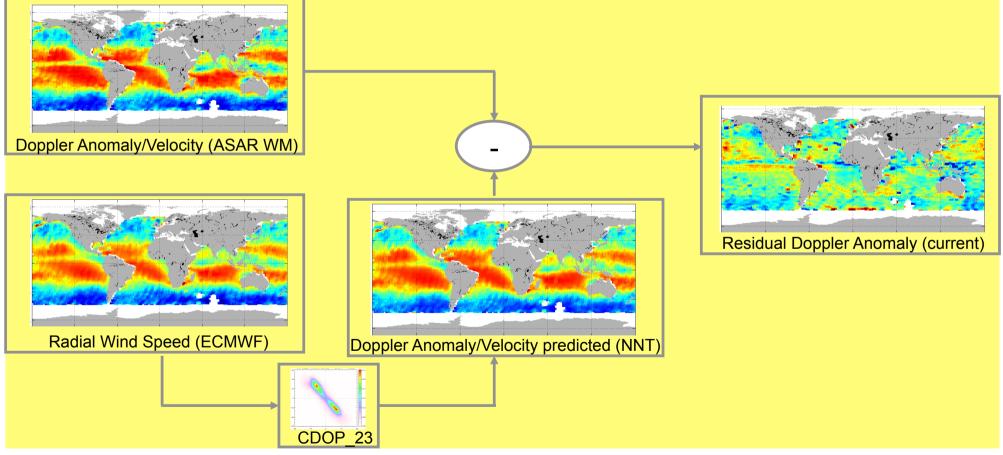
Simple methodology to remove wind effects





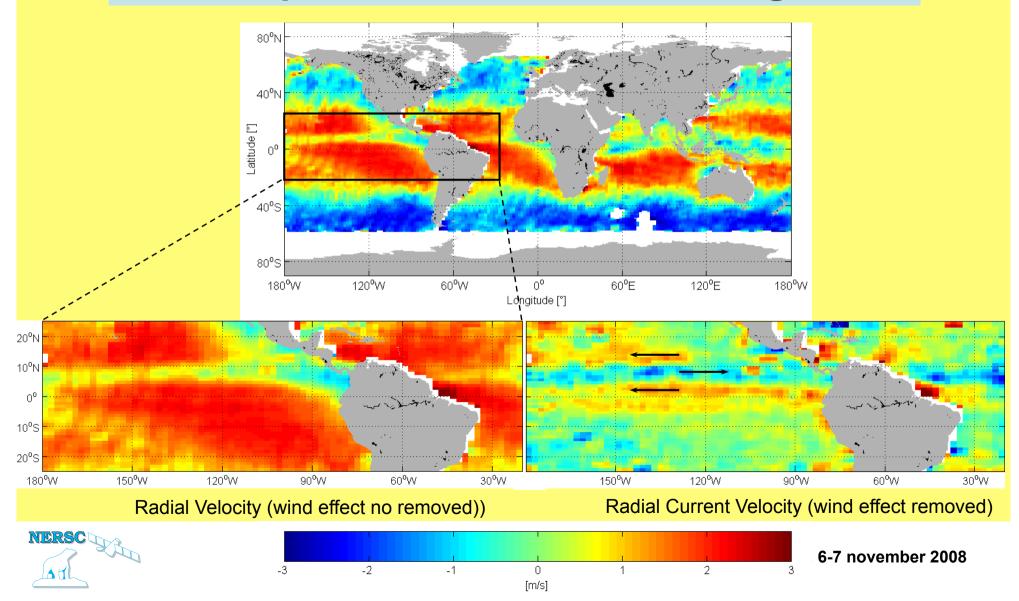
PART 3: THE NEW RESULTS

Simple methodology to remove wind effects

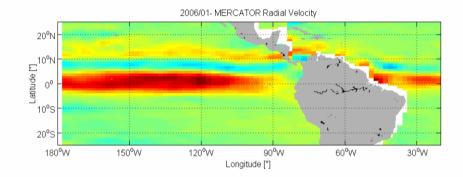


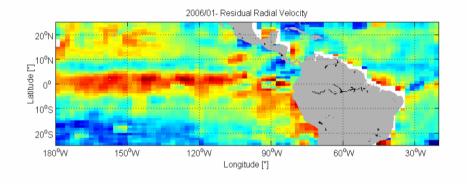


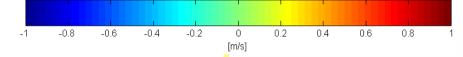
Equatorial Pacific monitoring

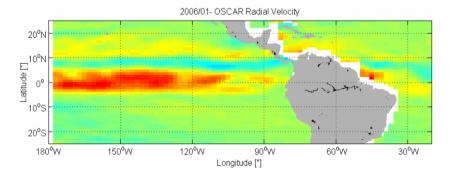


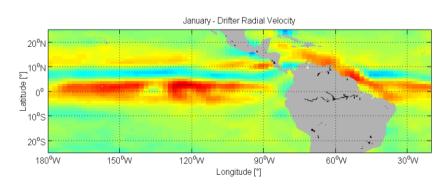
Equatorial Pacific monitoring: The seasonal cycle





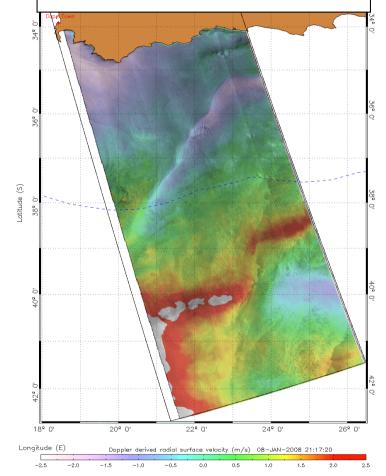




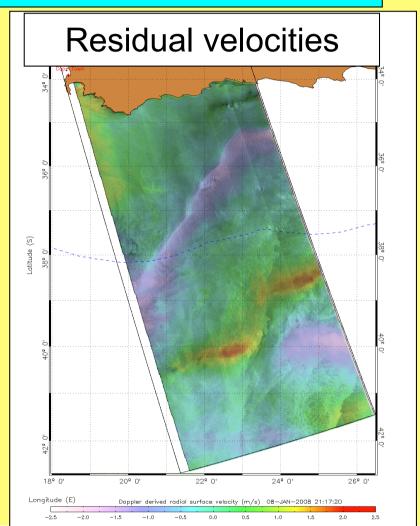


CDOP correction in Wide Swath

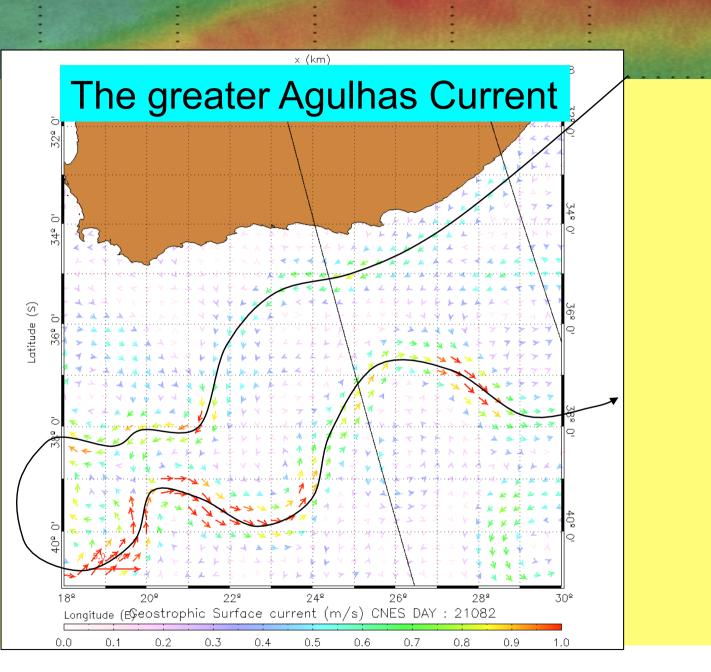
Total velocities



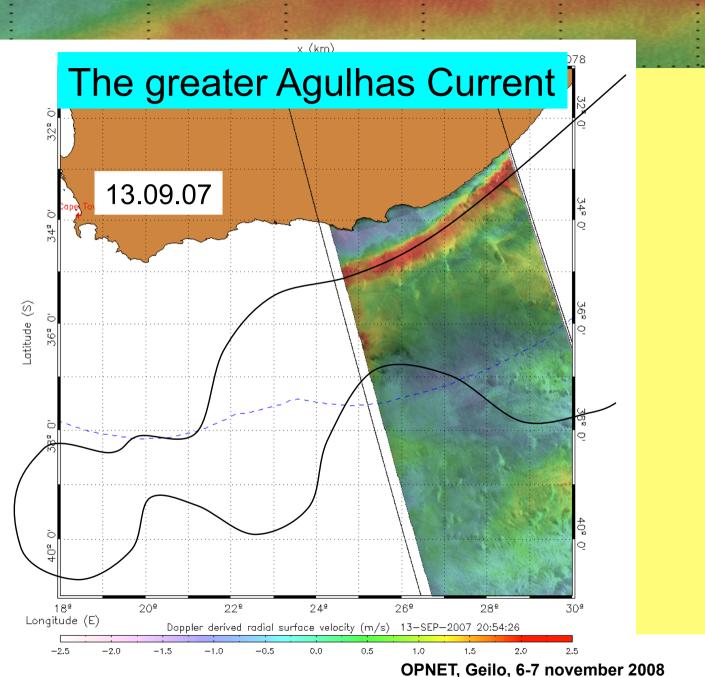
NERSC



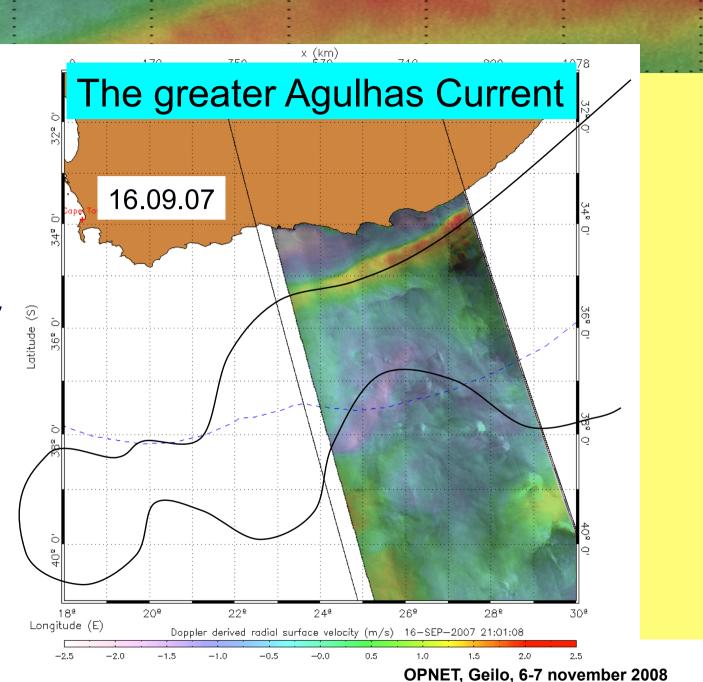




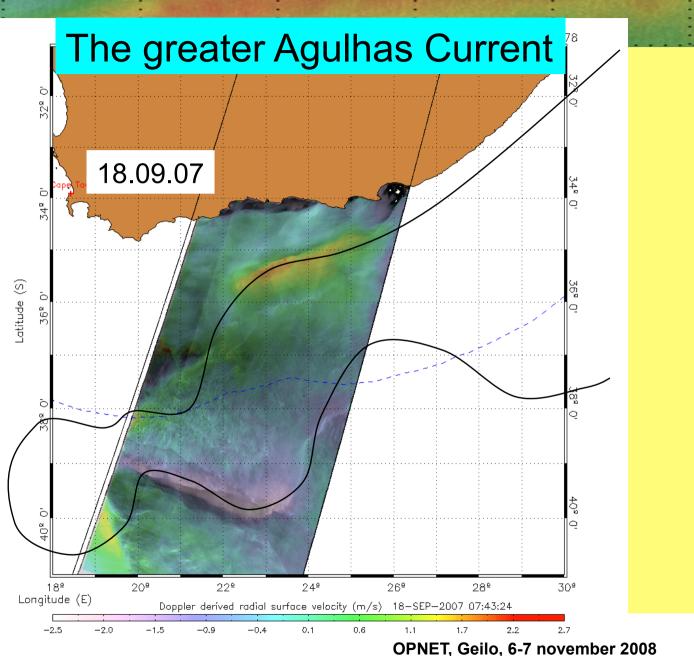
NERSC



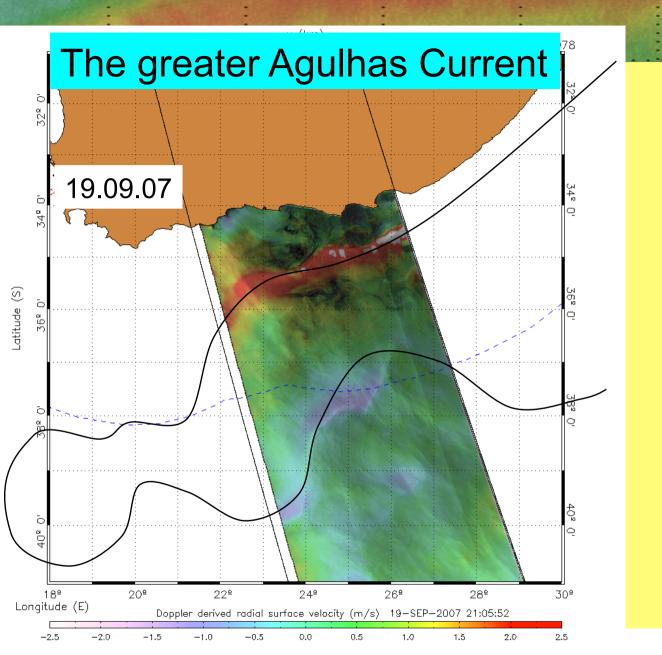




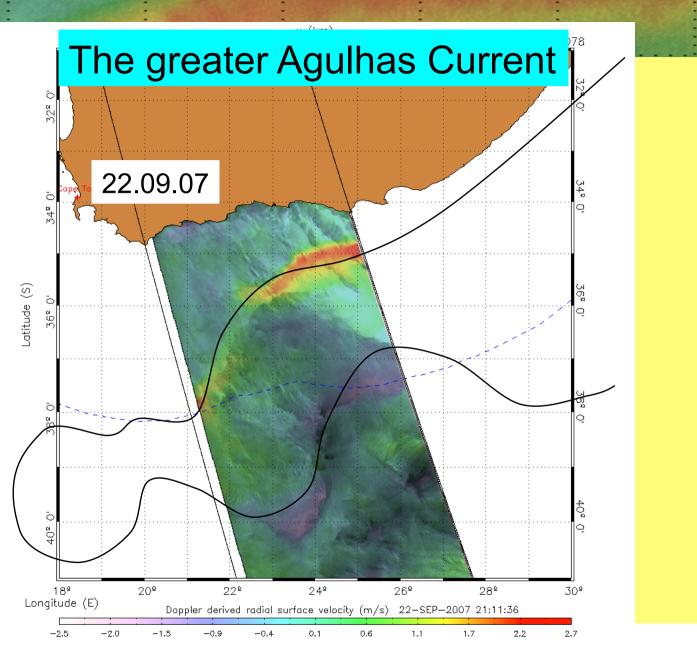






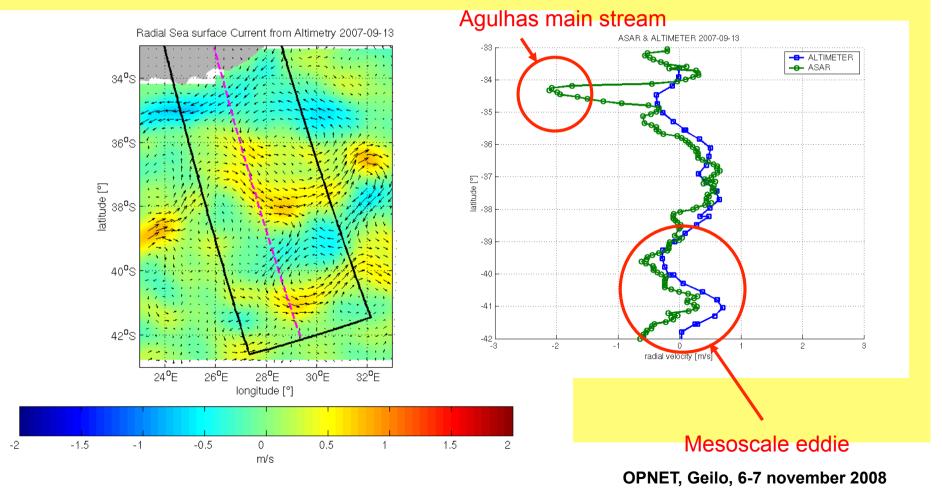






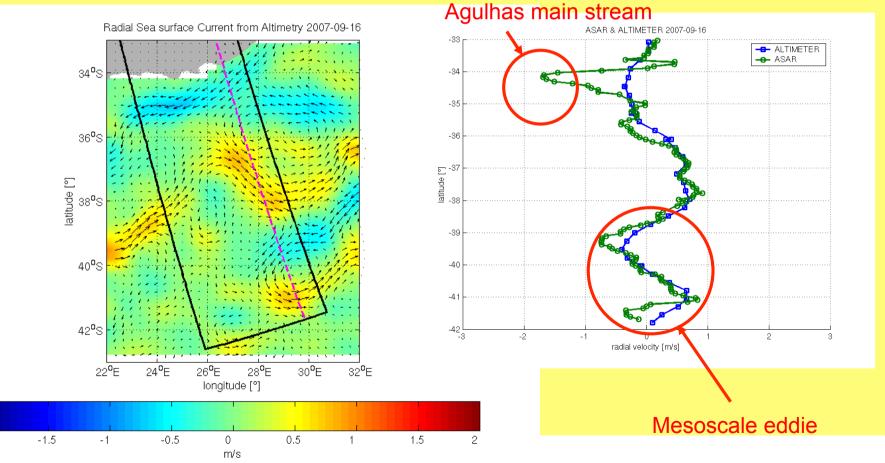


Validation exercise using altimetry



LATL

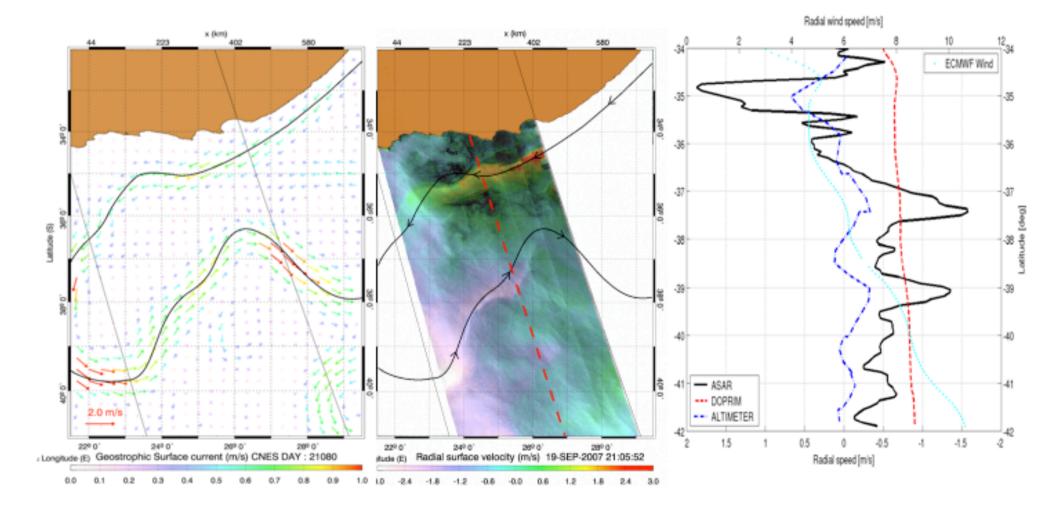
Validation exercise using altimetry



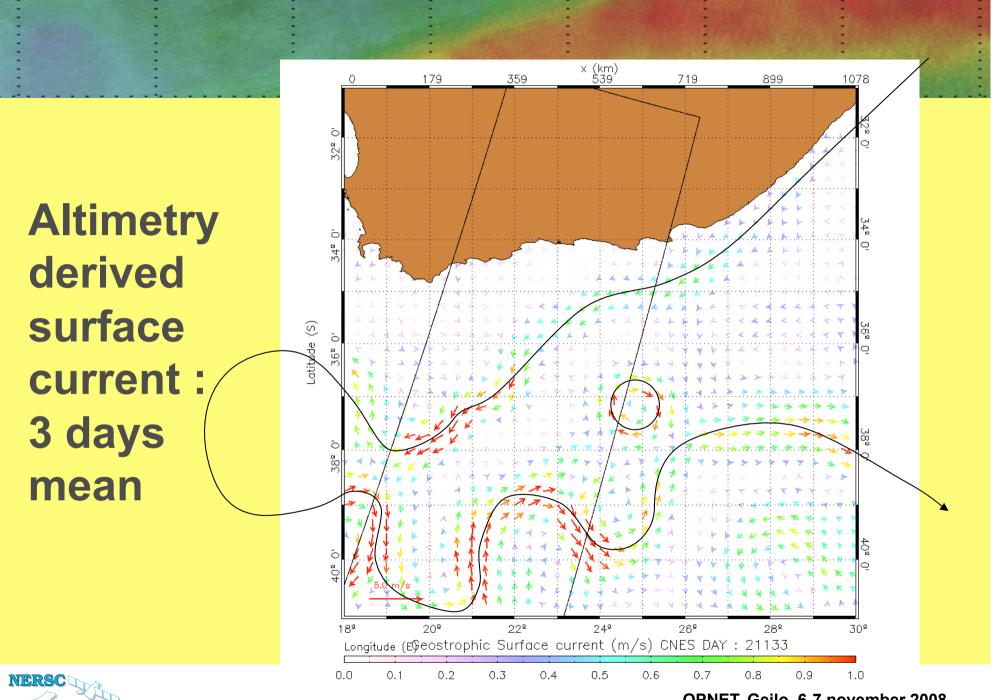
OPNET, Geilo, 6-7 november 2008

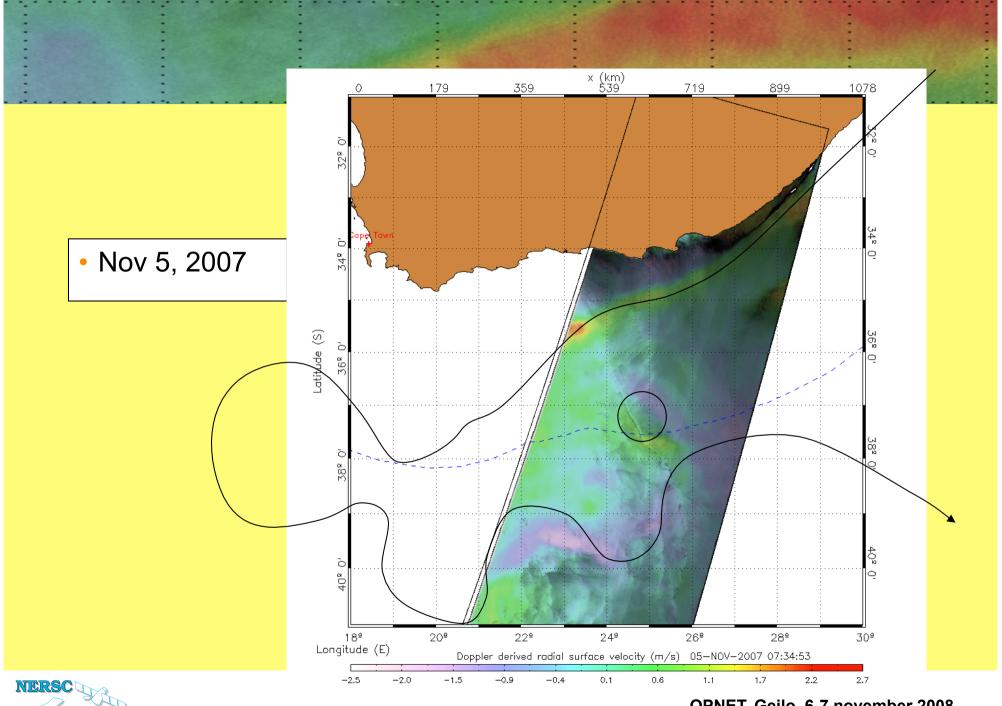
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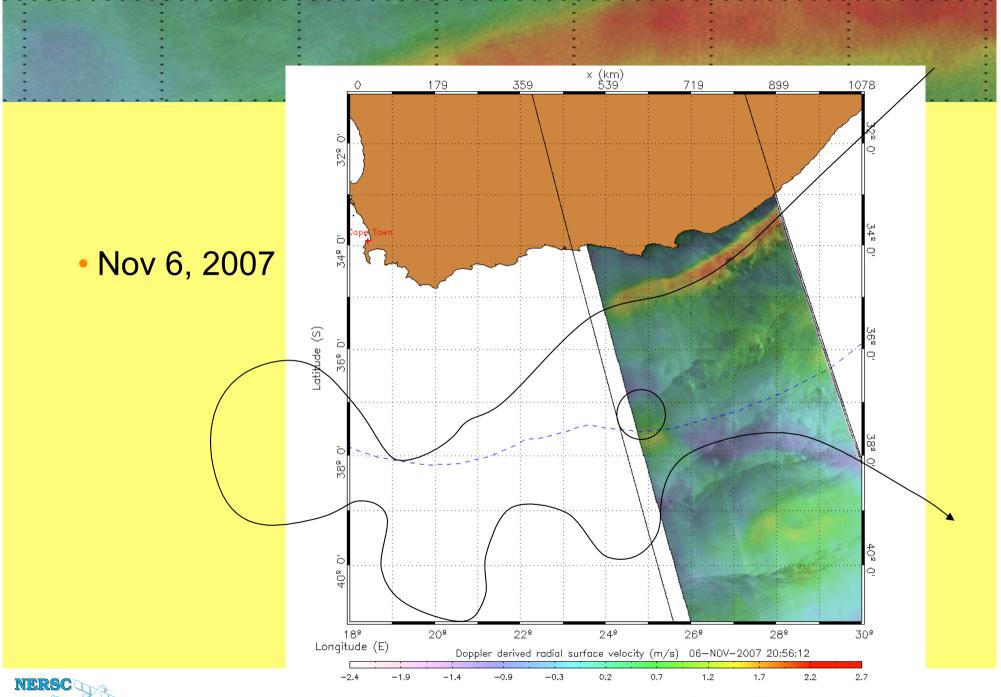
The greater Agulhas Current



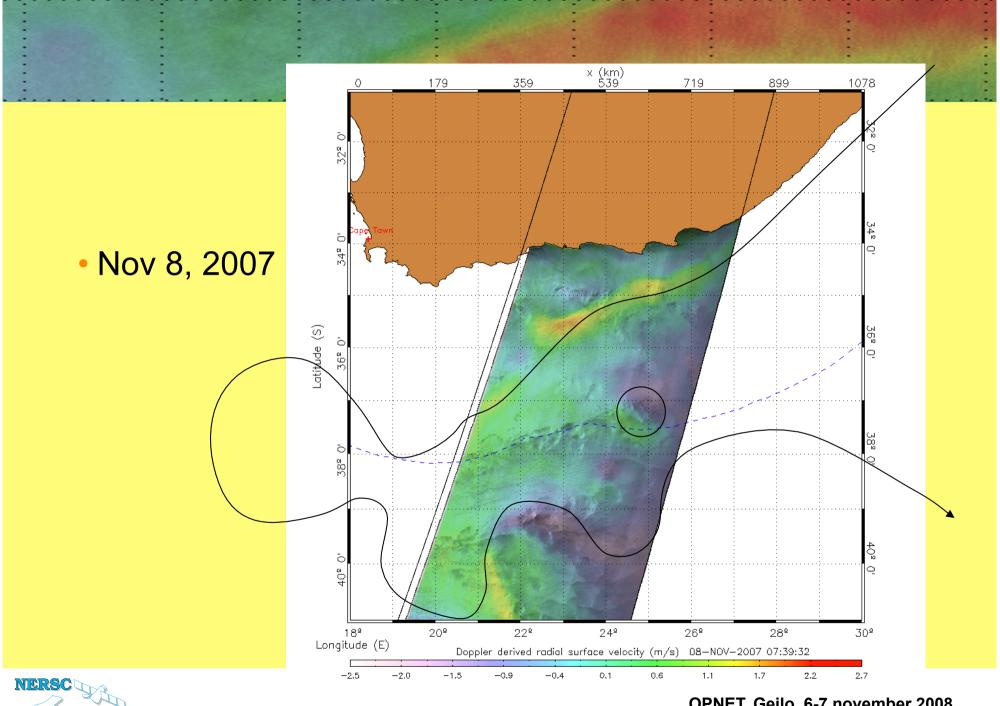


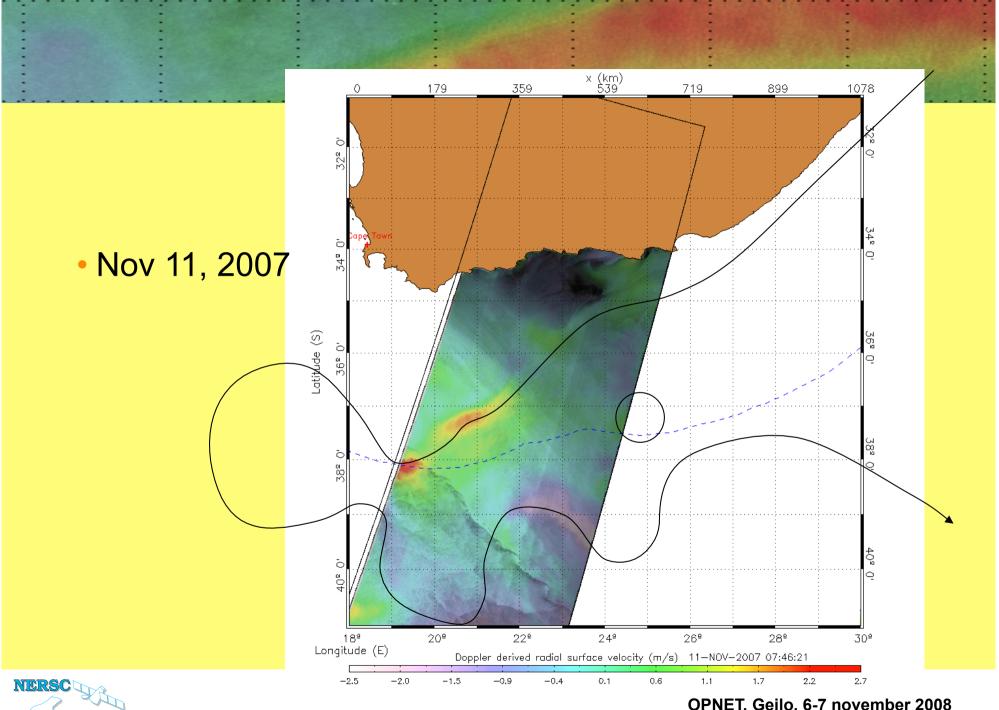


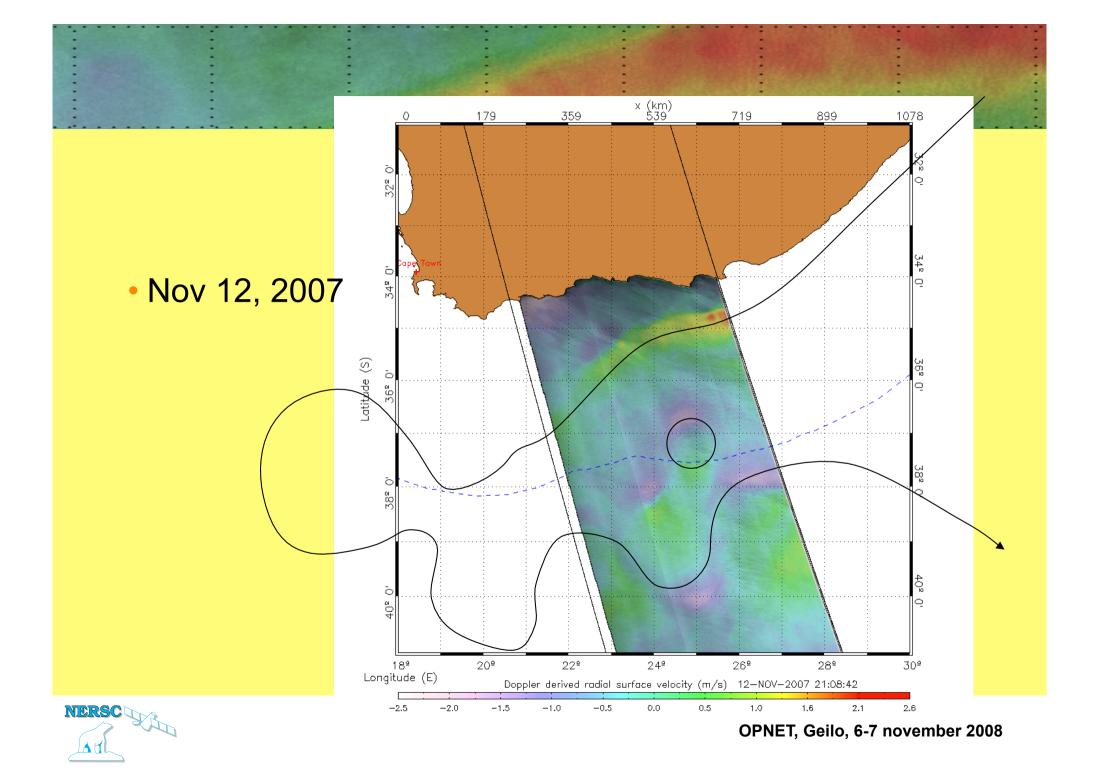


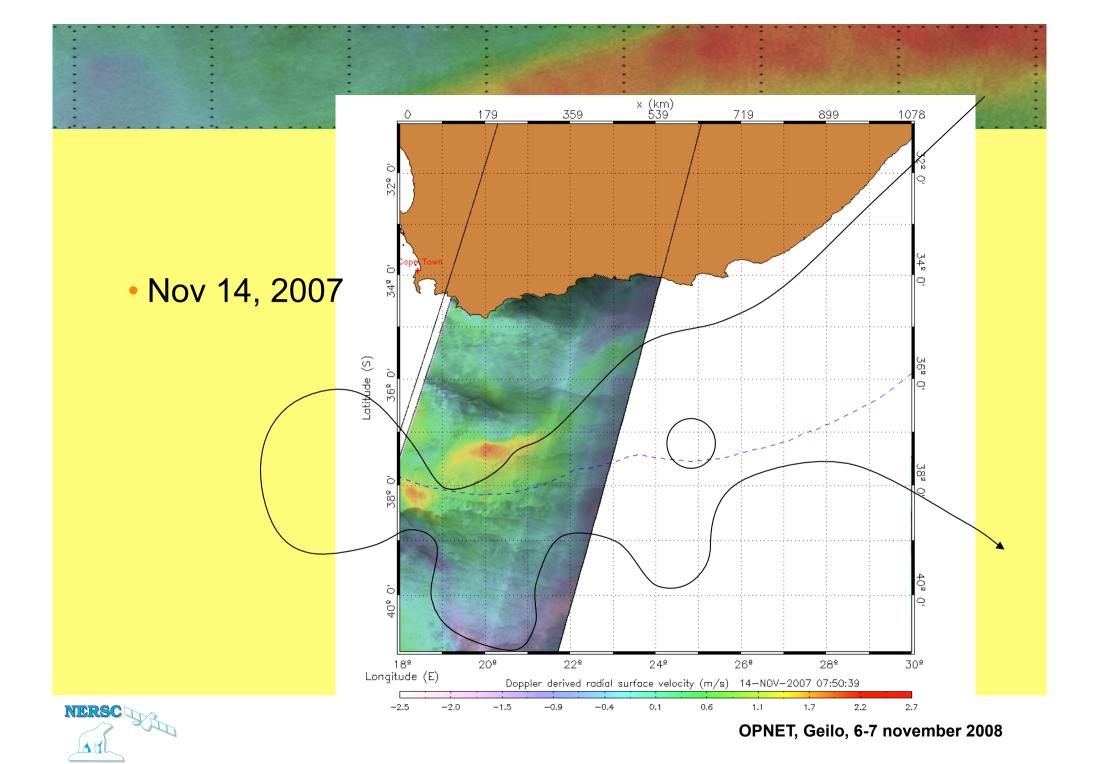


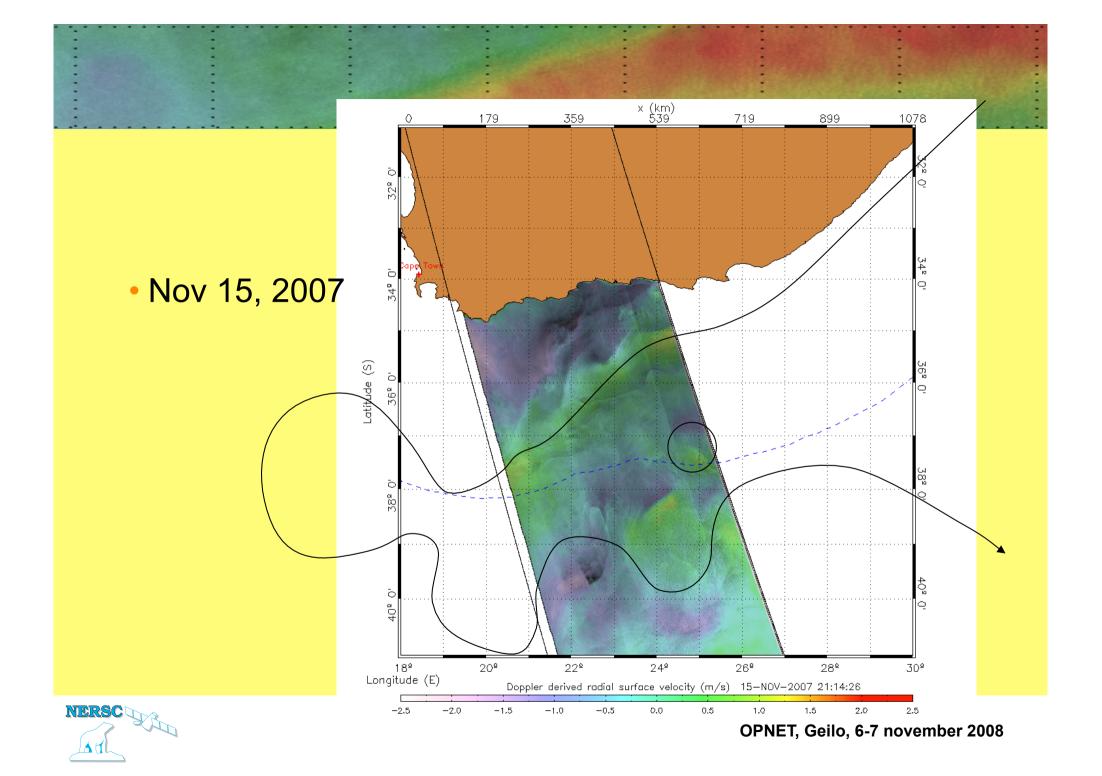
NERSC

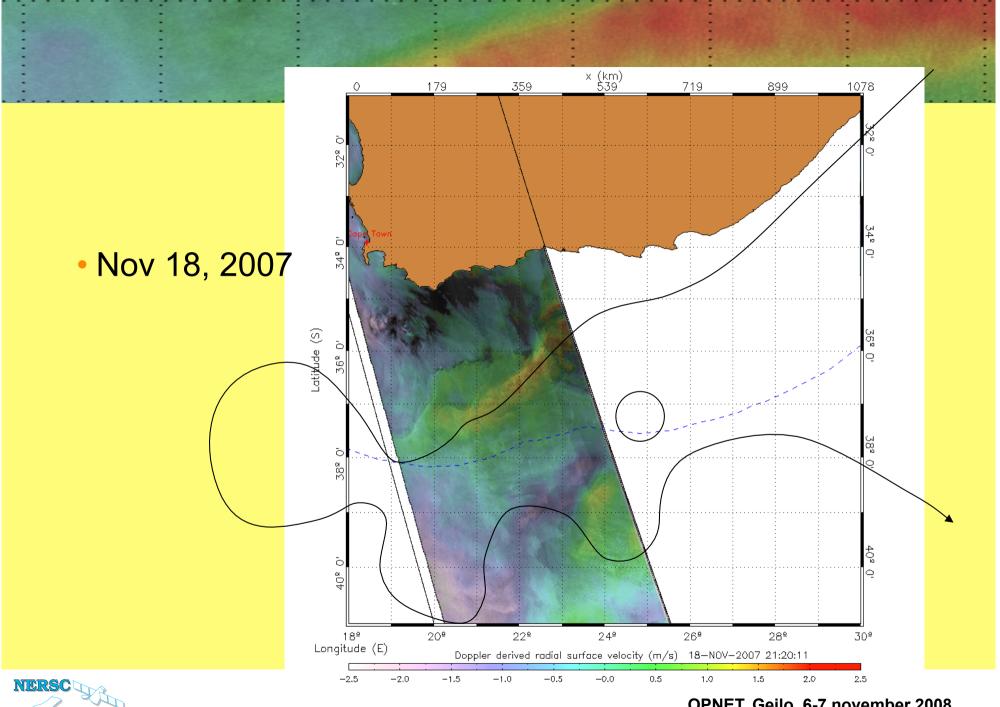


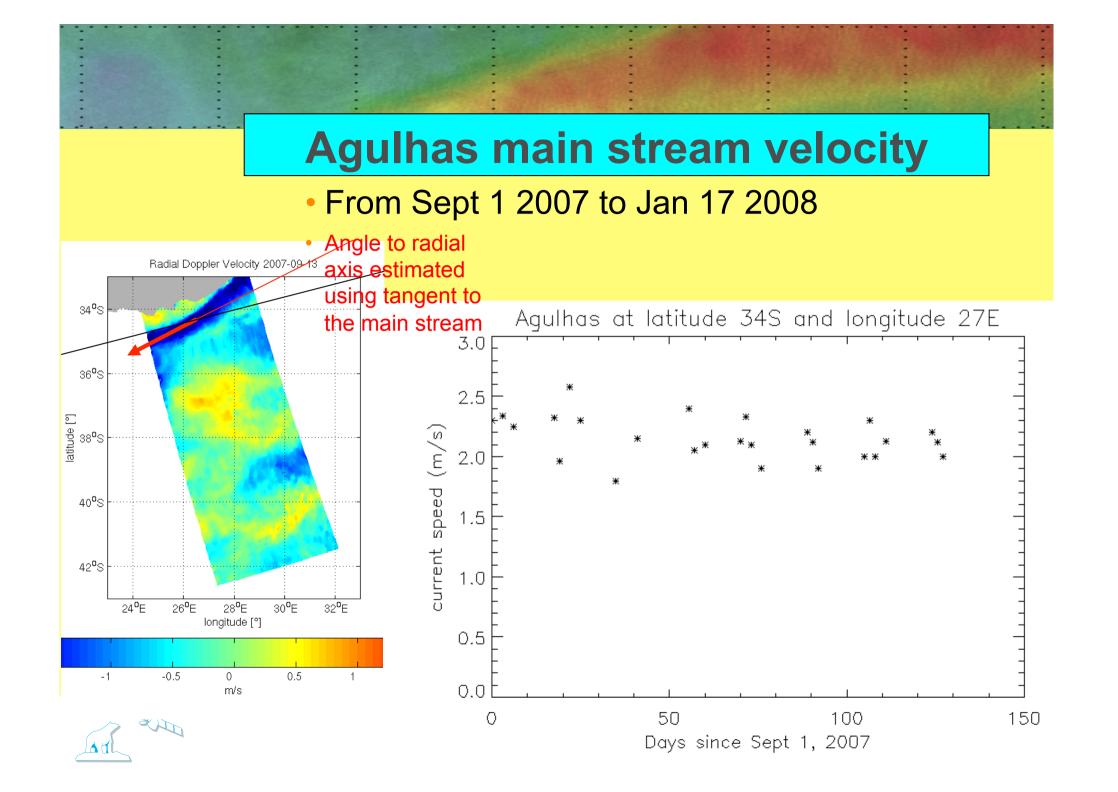










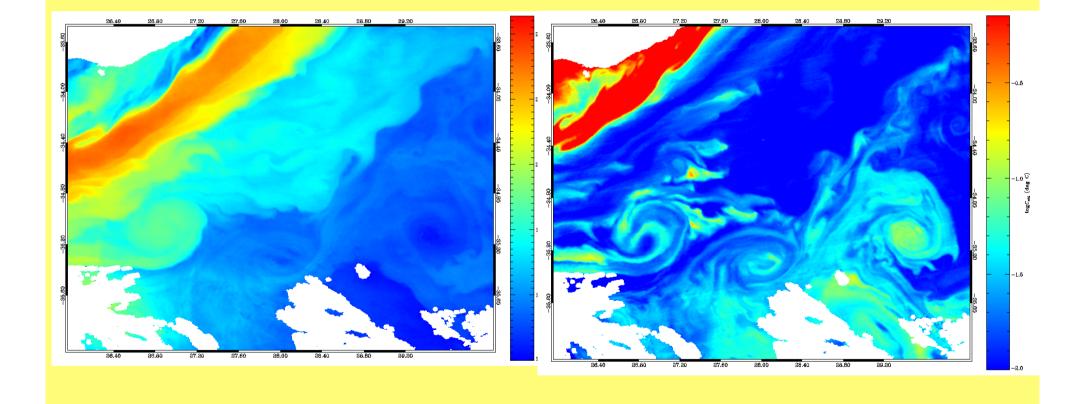


Summary

- Large scale surface velocity observed by SAR with possibility to obtain estimations of surface current in the radar look direction.
- CDOP applicable to the C band for removal of wind contribution to the Doppler velocity
- Combined with surface drifters and altimeter-derived surface geostrophic current, monitoring of the dynamics of intense current regimes may be advanced.
- Need for high resolution validation dataset. Routine acquisitions over HF radar is an option, but dedicated field campaign is needed.
- Accuracy and resolution of Doppler centroid estimation need more investigation.

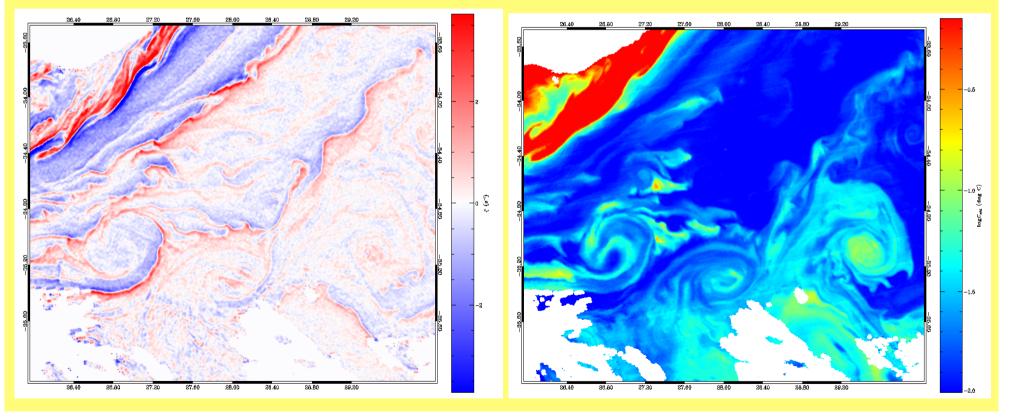


MODIS Brightness temperature and colour



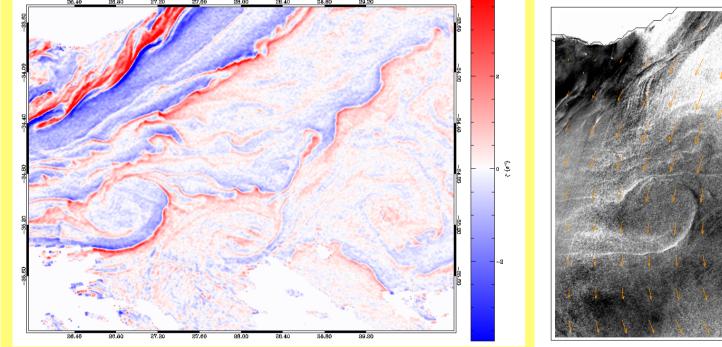


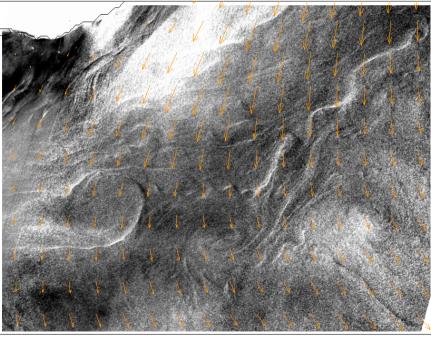
MODIS Brightness temperature SQG-derived vorticity $|\nabla \varphi|_{\varsigma=0}$ and colour





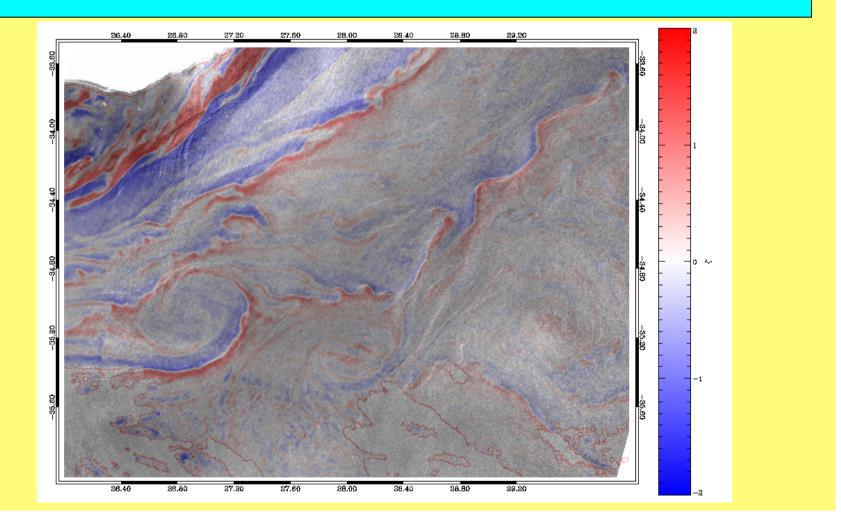
MODIS Brightness temperature SQG-derived vorticity $|\nabla \varphi|_{\varsigma=0}$ and ENVISAT radar roughness variations







Overlaid MODIS SQG-derived vorticity and ENVISAT radar roughness variations

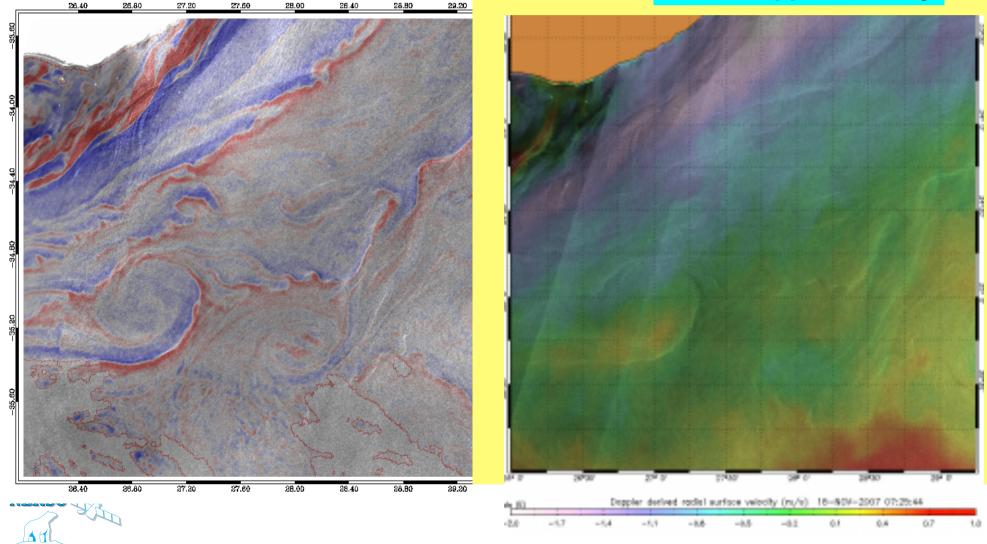


 $|\nabla \zeta|_{\zeta=0}$

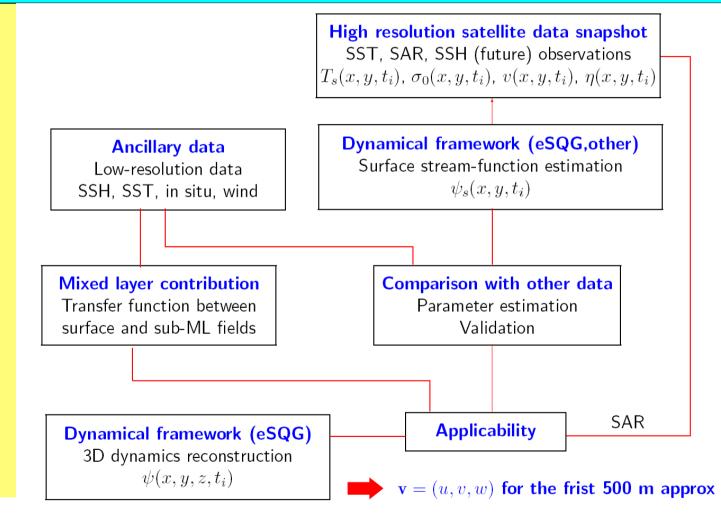


Overlaid MODIS SQG-derived vorticity and ENVISAT radar roughness variations

ASAR Doppler velocity



High resolution 3D ocean dynamics reconstruction from surface satellite data





The Doppler shift simulation in DopRIM

$$\frac{\pi f_D}{k_R} = -\frac{\overline{(u\sin\theta - w\cos\theta)\sigma_0(\theta + \Delta\theta)}}{\overline{\sigma_0(\theta + \Delta\theta)}}$$

$$f_{D}$$
: Doppler shift
 K_{R} : Radar Wave number
 θ : Incidence angle
U: horizontal velocity of scatterers
W: vertical velocity of surf scatterers
 σ_{0} : radar cross section

$$V_D = -\frac{\pi f_D}{k_R \sin \theta} = u_S + \overline{c}_f - \frac{1}{\tan \theta} \cdot \frac{\overline{\tilde{w}} \overline{\sigma}_0}{\overline{\sigma}_0} + \frac{\overline{\tilde{u}} \overline{\sigma}_0}{\overline{\sigma}_0}$$

 V_{D} : Doppler velocity C_{f} : Mean scatterer velocity Us: surface current in range σ_{0} : NRCS

 $V_D = u_S + \overline{c}_f + \overline{c}_f^{TH}$ $V_D = u_S + \sum P_f^p(\overline{c}_f + \overline{c}_f^{TH})$

f: br, sp, wb

p: pol

THIS IS THE BEAUTY OF DopRIM



